



Supplement of

Long-term observations of cloud condensation nuclei over the Amazon rain forest – Part 2: Variability and characteristics of biomass burning, long-range transport, and pristine rain forest aerosols

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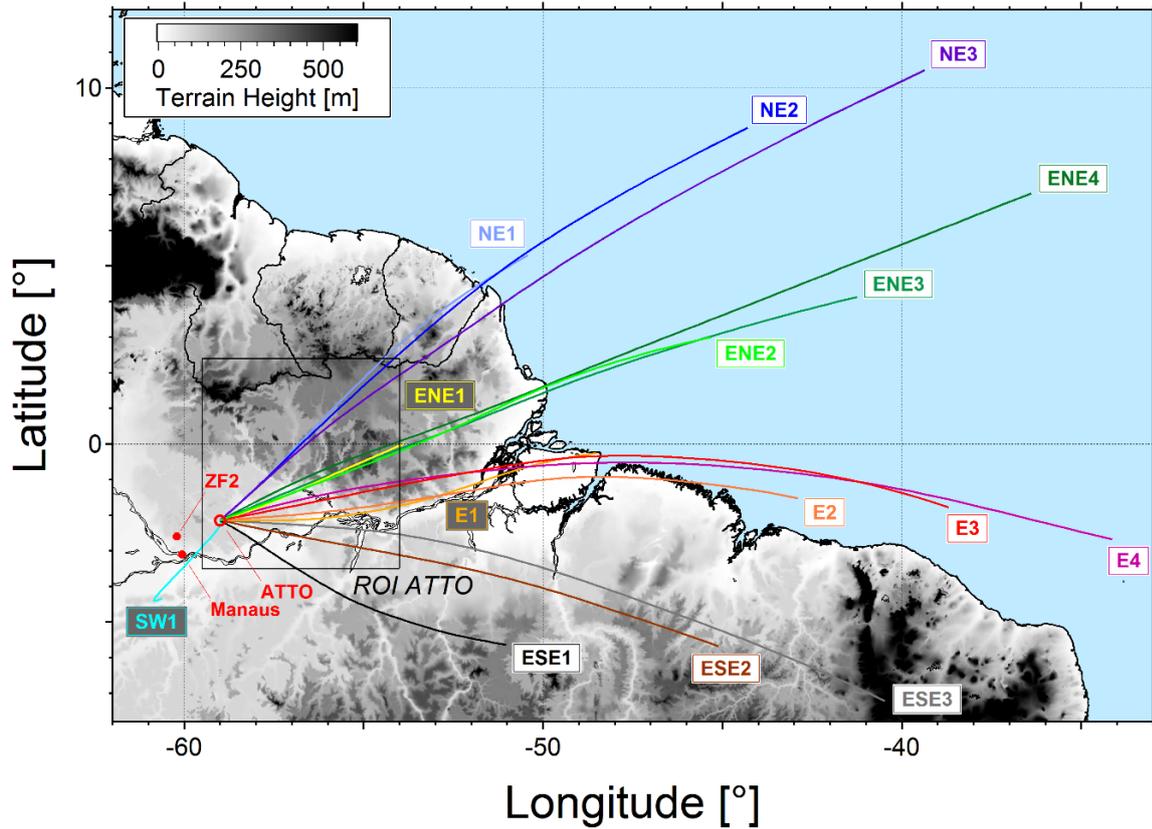


Figure S1. Map of northeast Amazon Basin with 15 clusters from *k*-means backward trajectory cluster analysis based on multiyear HYSPLIT data set (3-day BTs, 1 Jan 2008 - 30 Jun 2016, start height 1000 m AGL). Backward trajectory clusters show that air masses arrive at the ATTO site almost exclusively from northeastern to southeastern directions. Four major wind directions can be discriminated: (i) Northeastern (NE) clusters NE1, NE2, and NE3; (ii) east-northeastern (ENE) clusters ENE1, ENE2, ENE3, and ENE4; (iii) eastern (E) clusters E1, E2, E3, and E4; (iv) east-southeastern (ESE) clusters ESE1, ESE2, and ESE3. In addition, a separate cluster is directed towards the southwest (SW). A topographic map is represented by a grey scale, which is capped at 600 m. Highlighted area represents region of interest ROI_{ATTO} (59.5° W to 54° W and 3.5° S to 2.4° N). The figure has been adapted from C. Pöhlker et al. (2018), where a detailed description of the BT analysis can be found.

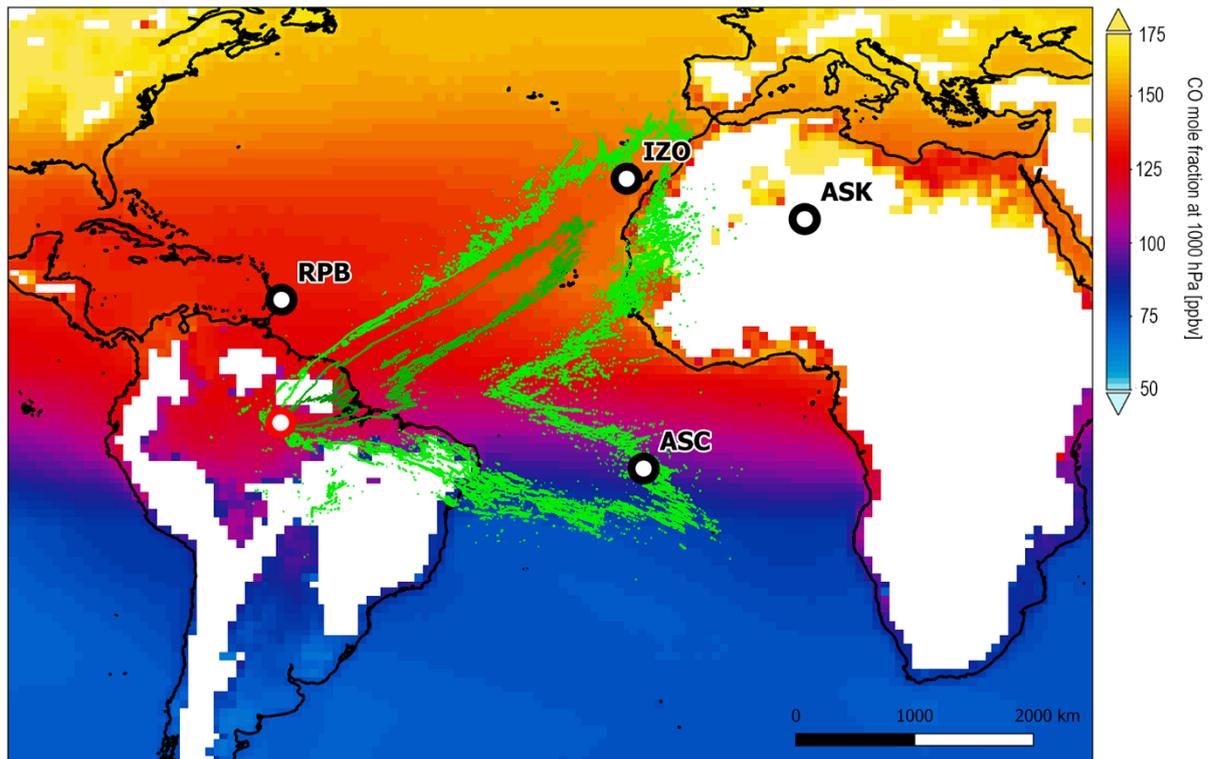


Figure S2. Satellite-derived map of CO mole fraction at ground level (1000 hPa, from AIRS instruments on board of the satellites Terra and Aqua) representing the average of the wet season periods (Feb to May) from 2012 to 2016. Green contour lines represent ATTO backward trajectory (BT) ensemble during wet season conditions (Feb to May) for years 2012 to 2016. Contour lines show upper 0.5 %, upper 5 %, and upper 25 % of BT-derived air mass residence times (for details see C. Pöhlker et al., 2018). Also shown are four background stations with monthly CO monitoring: Ragged Point, Barbados (RPB), Izaña, Tenerife, Canary Island (IZO), and Assekrem, Algeria (ASK) representing the northern hemispheric CO background vs. Ascension Island (ACS) representing southern hemispheric CO background. Data from CO background station was used in the definition of PR_{CO} filter for ATTO data (Sect. 2.7).

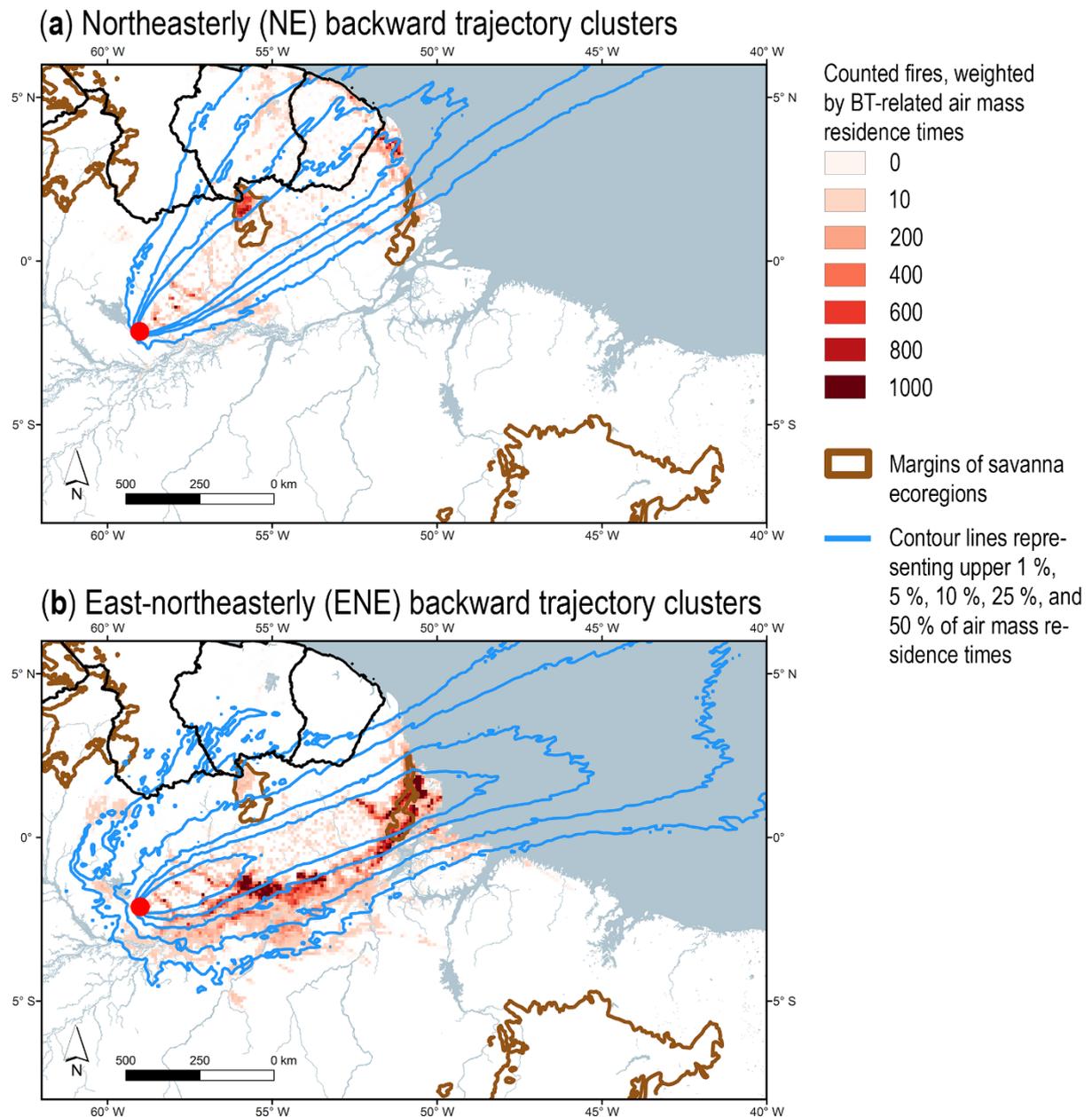


Figure S3. Fire maps being weighted with backward trajectory-derived air mass residence times for northeasterly BT clusters (i.e., NE1, NE2, NE3) and east-northeasterly BT clusters (ENE1, ENE2, ENE3, ENE4) according to C. Pöhlker et al. (2018). BT ensembles are shown as contour lines representing upper 1 %, 5 %, 10 %, 25 %, and 50 % of air mass residence times. Fire maps were created as outlined in C. Pöhlker et al. (2018). Geographic extent of savanna ecoregions has been obtained from Olson et al. (2001). Maps show that fire activity in footprint region of NE BT is overall rather low, with few fires in rain forest areas. However, relatively high fire abundance occurs in Guianan savanna region that is located in center of NE BT ensemble. This savanna region is still rather unperturbed by human activities and, thus, observed fires are likely predominantly wild fires according to de Carvalho and Mustin (2017). Fire activity in footprint of ENE BTs is mostly located along Amazon River valley, which is strongly influenced by human activities. Accordingly, these fire appear to be mostly man made.

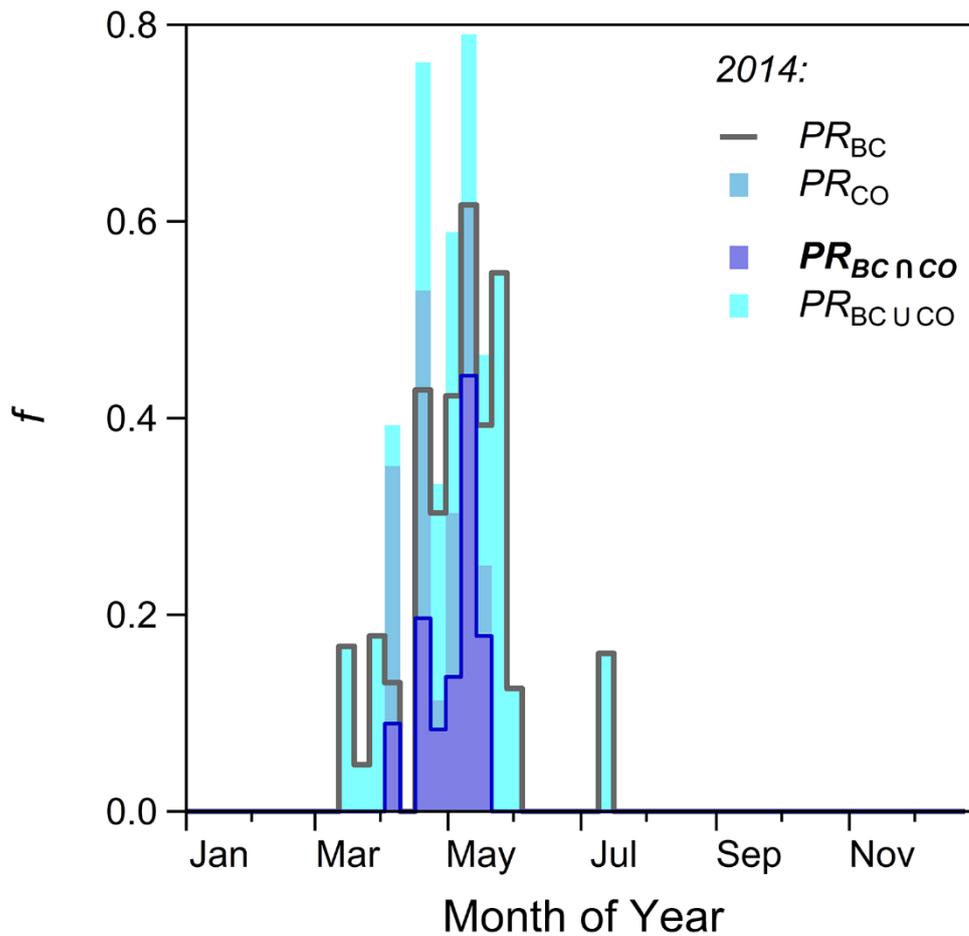


Figure S4. Seasonality of frequency of occurrence, f , of empirically pristine rain forest (PR) aerosol conditions at ATTO by means of PR_{BC} , PR_{CO} , $PR_{BC \cap CO}$, and $PR_{BC \cup CO}$ filters (see Sect. 2.7). $PR_{BC \cap CO}$ as most robust case is used as main PR filter throughout this work. Data is shown as weekly averages for year 2014, covering most of the CCN focal period of this study. For analogous representation of multi-year period, refer to Fig. 1.

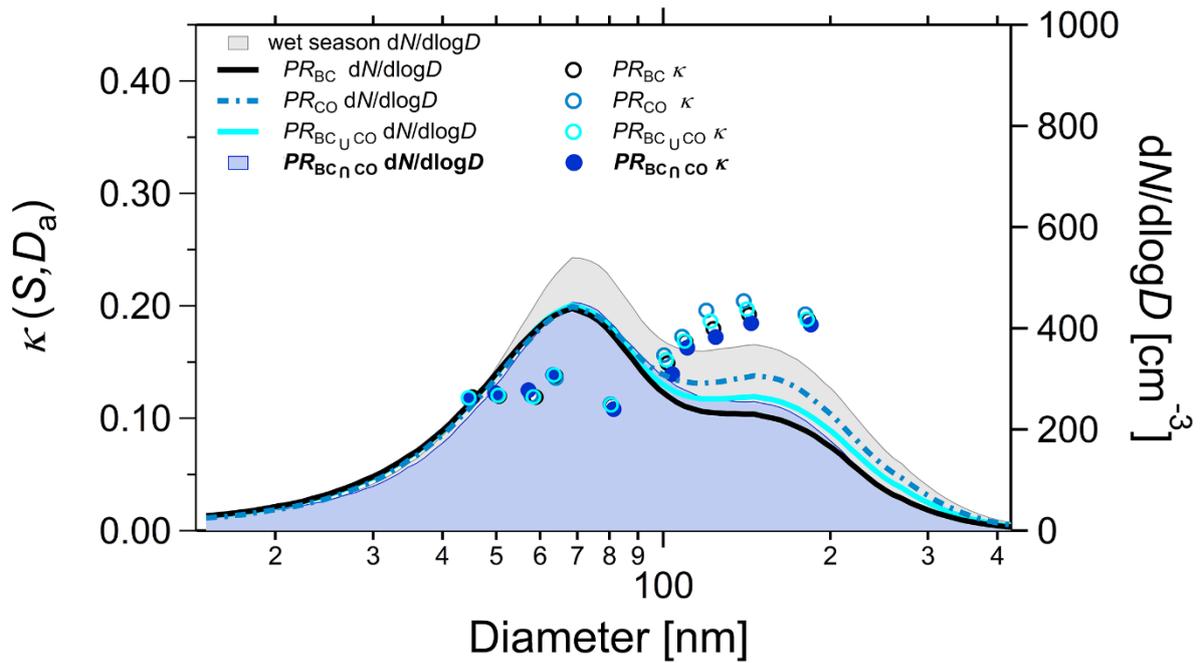
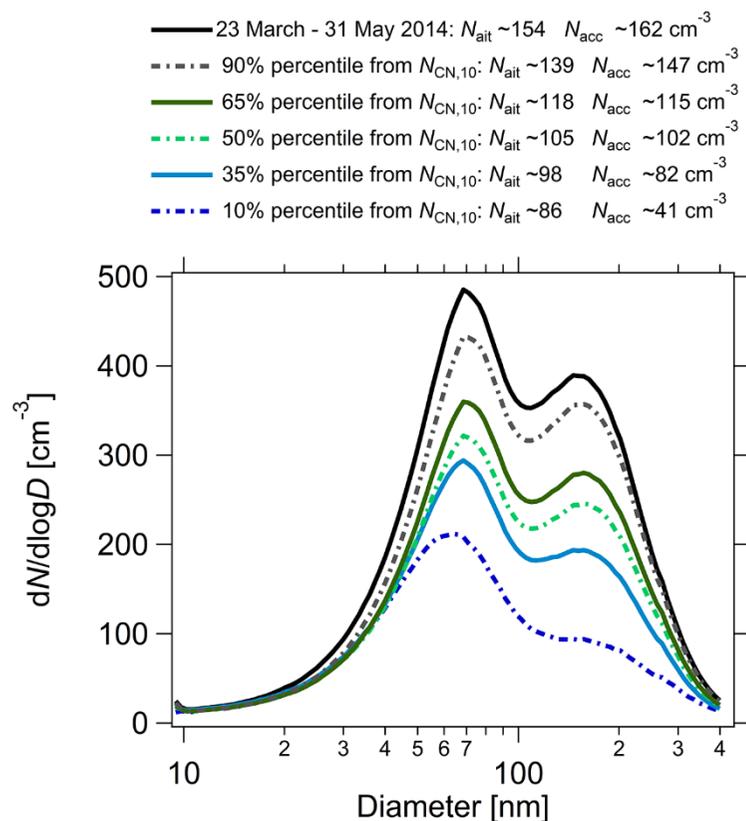


Figure S5. Size dependence of hygroscopicity parameter, $\kappa(S, D_a)$, and number size distributions of total aerosol particles, $N_{CN}(D)$, for all four *PR* filters (i.e., PR_{BC} , PR_{CO} , $PR_{BC \cap CO}$, $PR_{BC \cup CO}$, see Sect. 2.7). Values of $\kappa(S, D_a)$ for every S level are plotted against their corresponding midpoint activation diameter, $D_a(S)$. The average wet season $N_{CN}(D)$ size distributions (in grey) from the part 1 study (M. Pöhlker et al., 2016) has been added. The comparison shows that the *PR* filters give rather similar results, particularly in the Aitken mode size range. Certain, although minor, differences can be seen in the accumulation mode size range.

(a) Statistically filtered size distributions - wet season



(b) Statistically filtered size distributions - PR

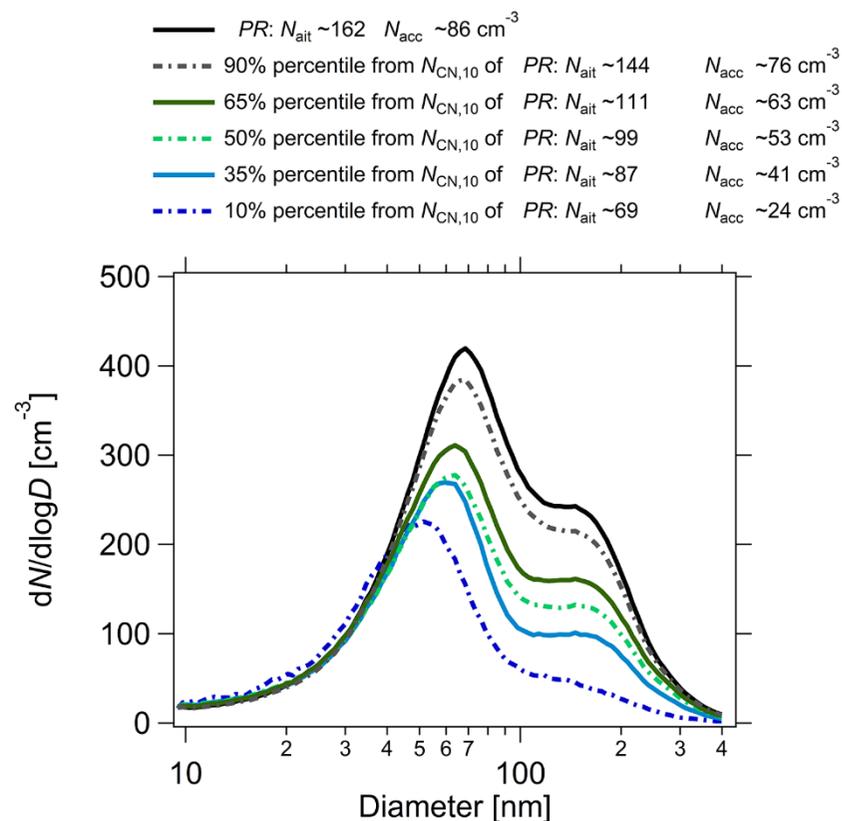


Figure S6. Number size distributions of total aerosol particles, $N_{CN}(D)$, classified by means of a statistical analysis of the total aerosol number concentration. Characteristic spectra for defined percentiles are shown for wet season vs. *PR* conditions. The comparison shows that the characteristic bimodal size distribution (i.e., relative strength of Aitken and accumulation modes) under wet season and *PR* conditions are rather stable. The distributions for the 10 % percentile of $N_{CN,10}$ deviate most clearly due to rain-related scavenging, which decreases the accumulation more strongly than the Aitken mode through activation of accumulation mode particles as CCN (compare Sect. 3.4).

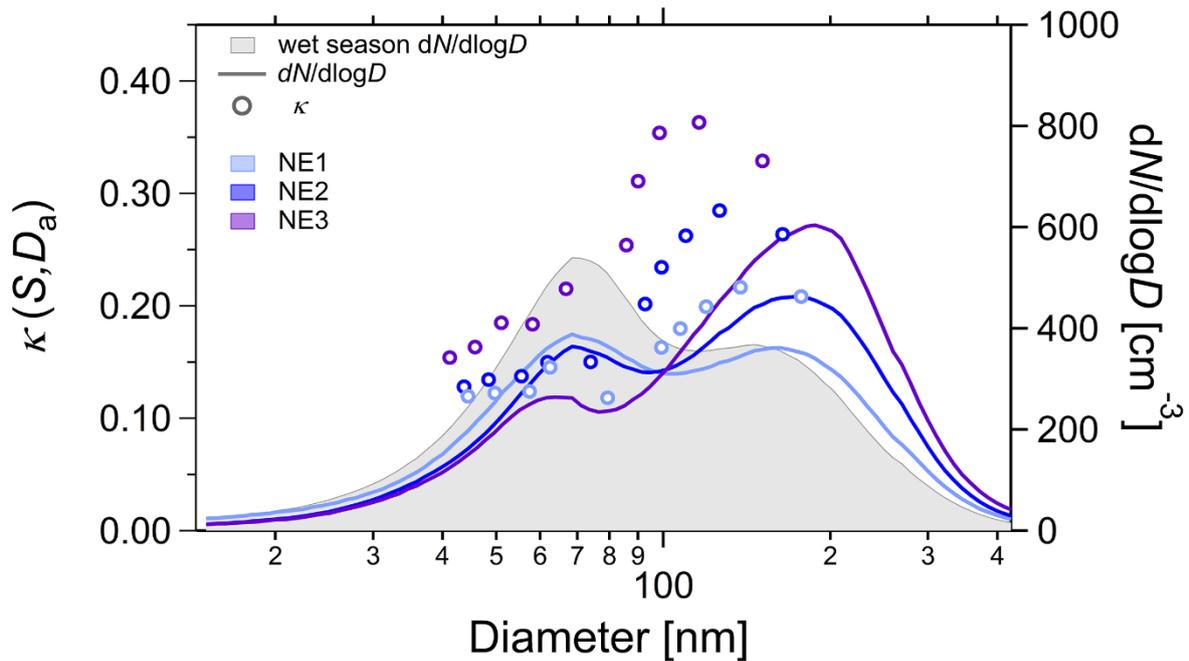


Figure S7. Size dependence of hygroscopicity parameter, $\kappa(S, D_a)$, and number size distributions of total aerosol particles, $N_{CN}(D)$, for northeasterly (NE) backward trajectory (BT) clusters according to C. Pöhlker et al. (2018). Values of $\kappa(S, D_a)$ for every S level are plotted against their corresponding midpoint activation diameter, $D_a(S)$. The average wet season spectrum from the companion part 1 study (M. Pöhlker et al., 2016) has been added for comparison. The figure shows rather pronounced differences between the three NE clusters due to the fact that the NE3 BTs are most efficient in transporting African LRT plumes into the Amazon Basin during its wet season (compare Moran-Zuloaga et al., 2017). The NE1 BTs are least efficient in LRT transport due to strong rain-related scavenging. The NE2 BTs represent an intermediate case. Accordingly the $N_{CN}(D)$ and $\kappa(S, D_a)$ size distributions for NE3 resemble the LRT results in Fig. 6c, the NE1 case is most similar to the PR results in Fig. 6a, whereas the NE2 case resembles a mixture of both.

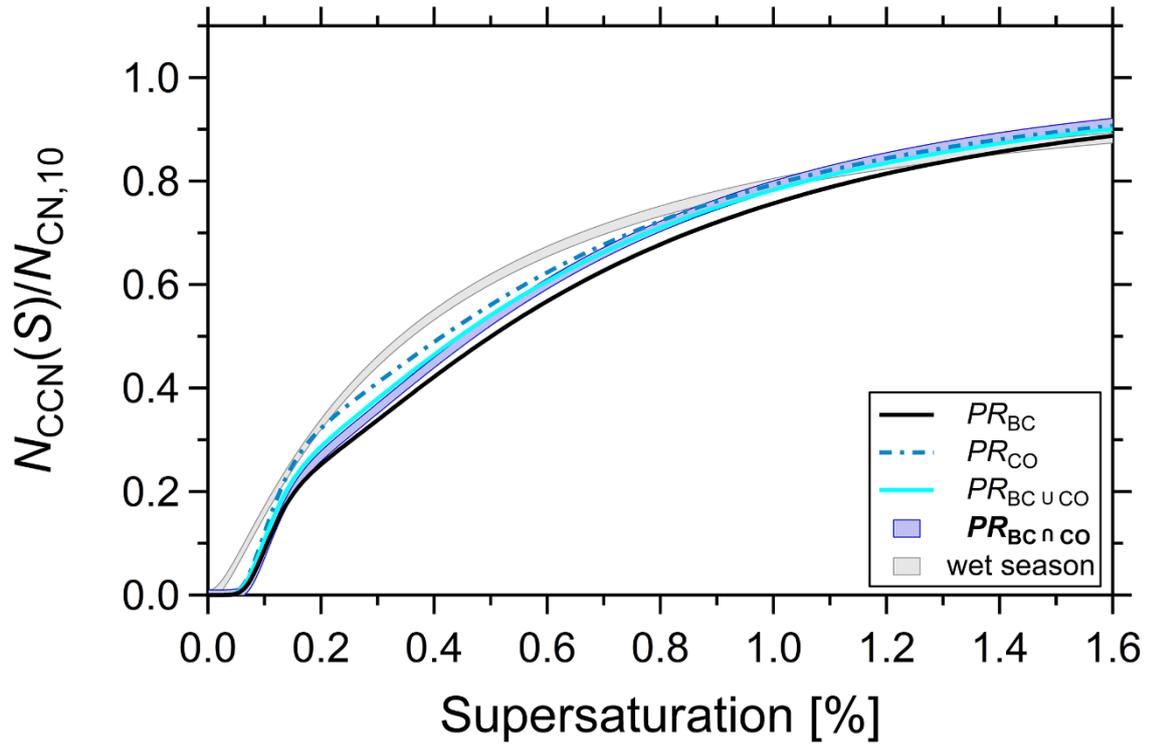


Figure S8. Erf fits of CCN efficiency spectra for all PR filters (i.e., PR_{BC} , PR_{CO} , $PR_{BC \cap CO}$, $PR_{BC \cup CO}$, see Sect. 2.7). The average CCN efficiency spectrum for wet season conditions from companion part 1 study (M. Pöhlker et al., 2016) has been added for comparison. The figure shows that the PR filters give rather similar CCN efficiency spectra. Minor differences in accumulation mode strength as shown in Fig. S5 result here in slightly different shapes of PR CCN efficiency spectra.

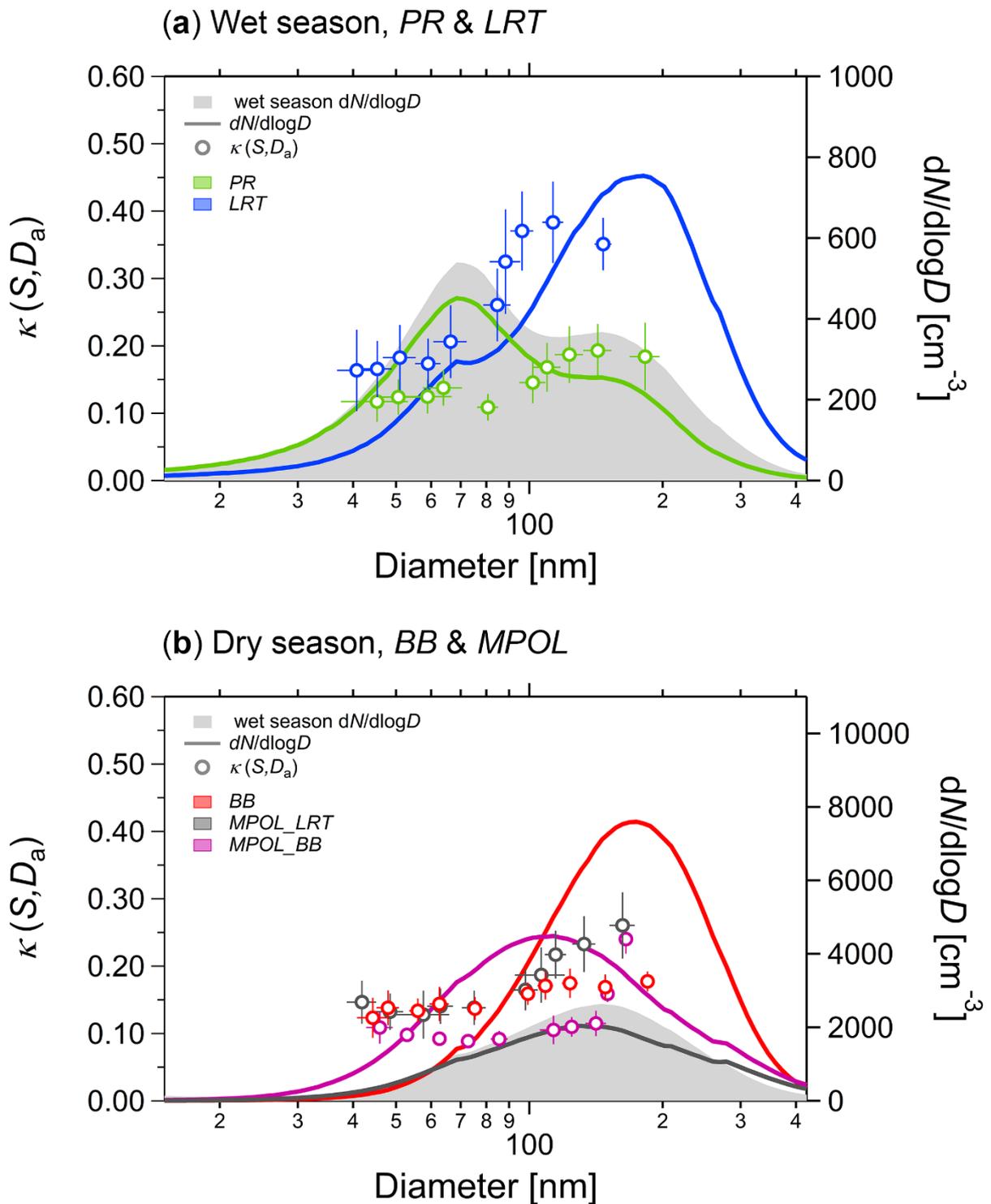


Figure S9. Overview figure combining $N_{CN}(D)$ and $\kappa(S, D_a)$ size distributions for *PR*, *LRT*, *BB*, *MPOL-LRT*, and *MPOL-BB* conditions as well as average wet and dry season results from companion part 1 paper (M. Pöhlker et al., 2016). Values of $\kappa(S, D_a)$ for every S level are plotted against their corresponding midpoint activation diameter $D_a(S)$. For $\kappa(S, D_a)$, the error bars represent one std. For $D_a(S)$, the experimentally derived error is shown.

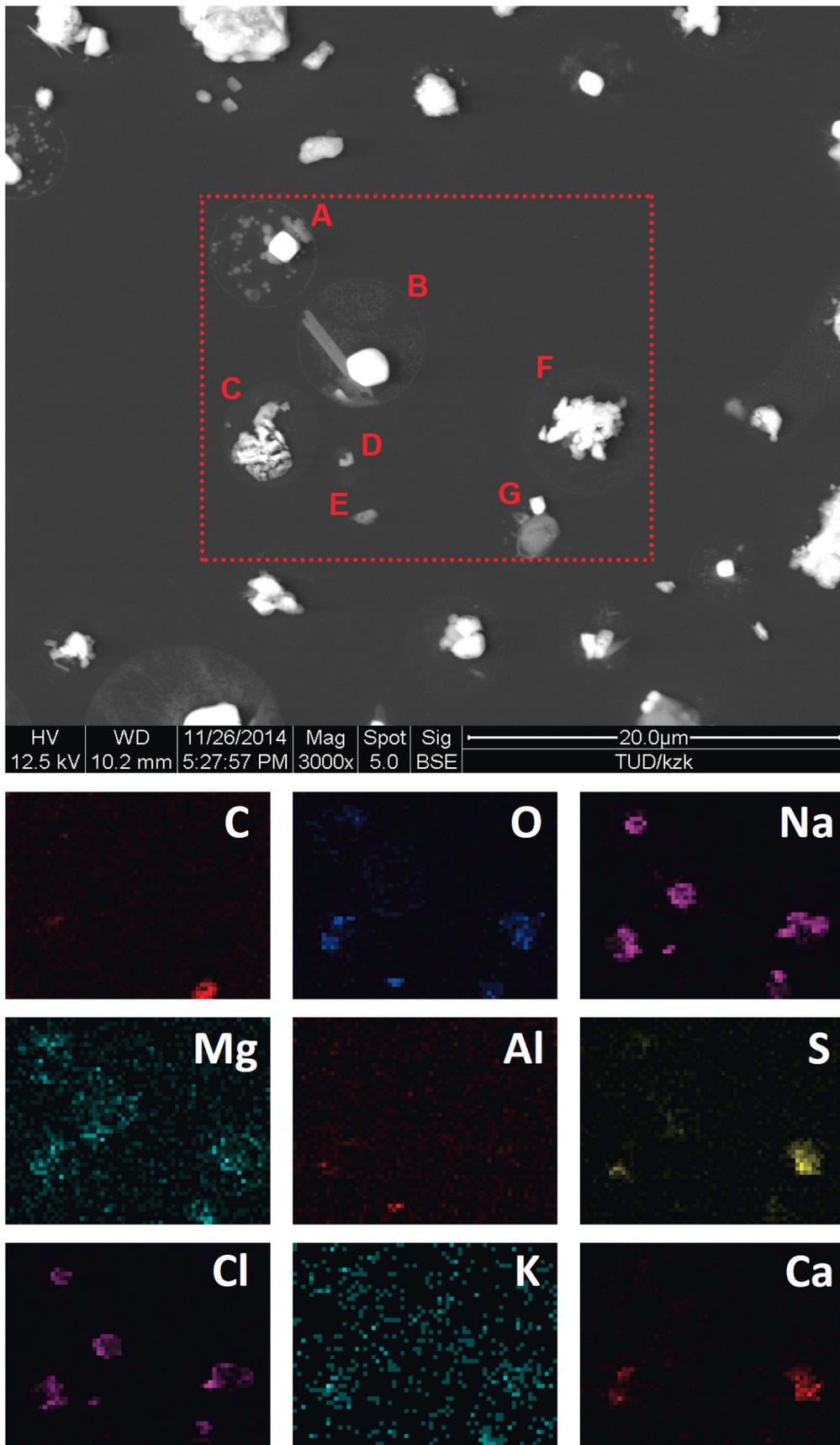


Figure S10. SEM images and EDX maps for selected aerosol sample showing typical appearance of LRT aerosol population at ATTO. Sample shows processed and internally mixed sea spray and dust particles. The results are based on an aerosol sample, collected on 15th Feb 2014 during LRT event 2014_2 according to Moran-Zuloaga et al. (2017), which is comparable (i.e., transport patterns,

physicochemical properties recorded by online measurements) to the event discussed as LRT case study in the present work. The particles A, B, and D resemble sea spray particles with a (large) cubic crystal of NaCl as well as weaker signals in Mg, O, and S (likely as MgSO₄). Morphologically, the particles show similarities to examples of aged sea-salt particles in previous studies (e.g., Laskin et al., 2012). The needle-like structure in particle B could be a CaSO₄ crystal that appears to be too 'thin' to give a clear signal in the EDX maps. The particles C and F also show a strong NaCl signal. In addition, they reveal strong signals in S, O, and Ca. These particles resemble CaCO₃ particles from the Saharan dust plume, in which the CO₃²⁻ anion was (partly) replaced by SO₄²⁻ in the course of atmospheric processing. Similar particles have been described in Laskin et al. (2005). Upon processing, a certain amount of NaCl appears to be mixed in most particles. Particle G comprises a NaCl crystal that is attached to PBAP-like particle (based on morphology and strong C signal). Particle E shows signals in Al and O, which fit to an aluminosilicate particle.

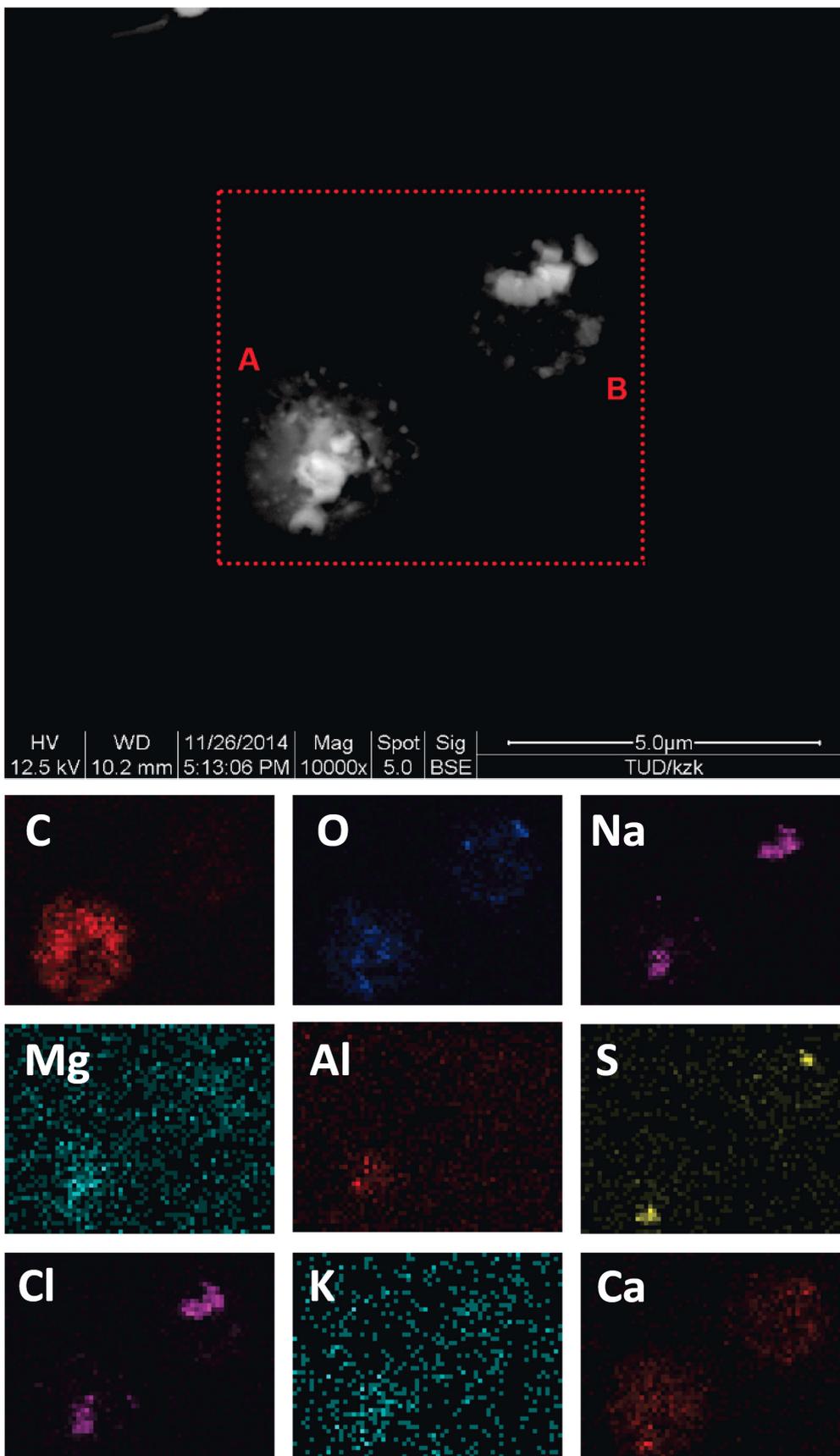


Figure S11. SEM images and EDX maps for selected aerosol sample showing a further example of the typical appearance of LRT aerosol population at the ATTO site. Sample shows processed and internally mixed sea spray and dust particles. For further relevant details, refer to caption on Fig. S10.

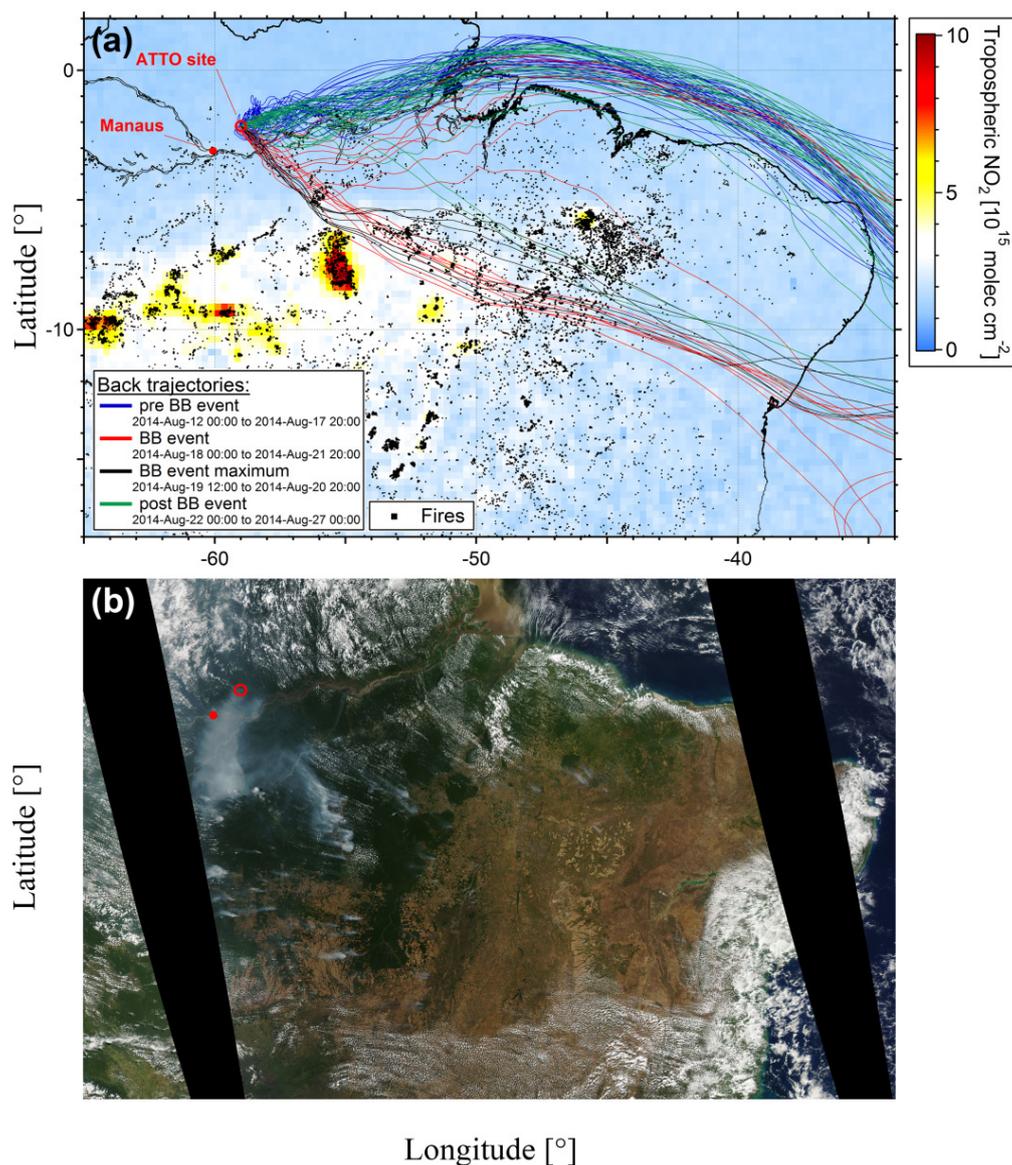


Figure S12. Composite maps illustrating origin of strong biomass burning (BB) plume that affected the ATTO site during *BB* case study (17 to 23 Aug 2014). (a) Map combines BT ensembles for selected periods [HYSPLIT, NOAA-ARL, GDAS1, start height 200 m, BTs started every 4 h, Draxler and Hess (1998)] with satellite products for active fire counts [MODIS active fire detection extracted from MCD14ML distributed by NASA FIRMS, Justice et al. (2011)] and NO₂ atmospheric column (OMI/Aura NO₂ Cloud-Screened Total and Tropospheric Column Daily L3 Global 0.25deg) during same period. BTs are color-coded for characteristic periods before, during, and after biomass burning event. BTs for pollution ‘peak’ at ATTO are color-coded as “BB event maximum”. Map indicates that the BTs’ main track temporarily swings from ‘coastal path’ to ‘inland path’ and back. Inland BTs intersect areas with strong fire activity, as visible in fire counts as well as in NO₂ map (i.e., 7° S, 55° W). The most intense fires are located along the highway BR-163, which is known to be a hotspot of recent deforestation activities – for details, see C. Pöhlker et al. (2018). (b) MODIS corrected reflectance image (taken during Aqua overflight on 19 August 2014 1700UT) confirms that major fires at about 7° S and 55° W have emitted clearly visible smoke plume that travelled northwestwards and impacted the ATTO site during the *BB* case study.

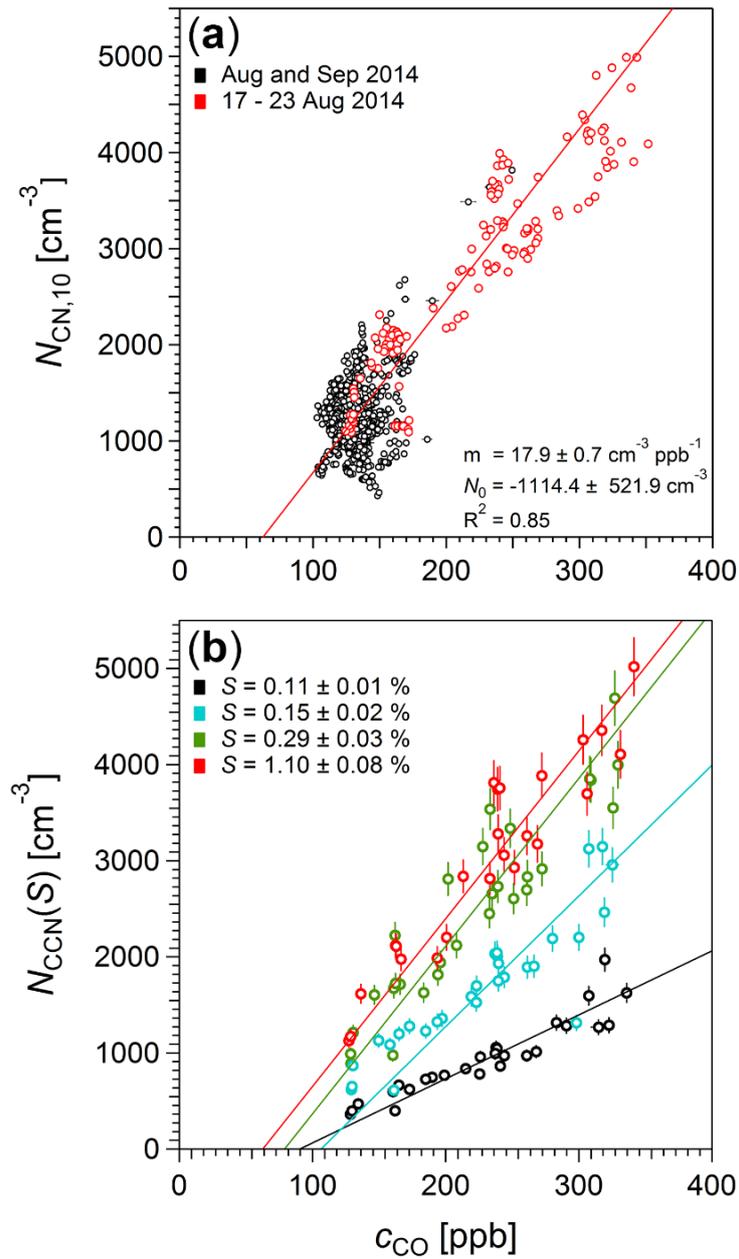


Figure S13. Scatter plots of total aerosol concentration, $N_{CN,10}$, and CO mixing ratio, c_{CO} , as hourly average values with bivariate regression fit in (a) and scatter plot of CCN concentrations, $N_{CCN}(S)$, and CO mixing ratio, c_{CO} , for four selected S with bivariate regression fit in (b). Data in both cases covers peak period of the biomass burning case study *BB* (see Fig. 4). Error bars show standard errors.

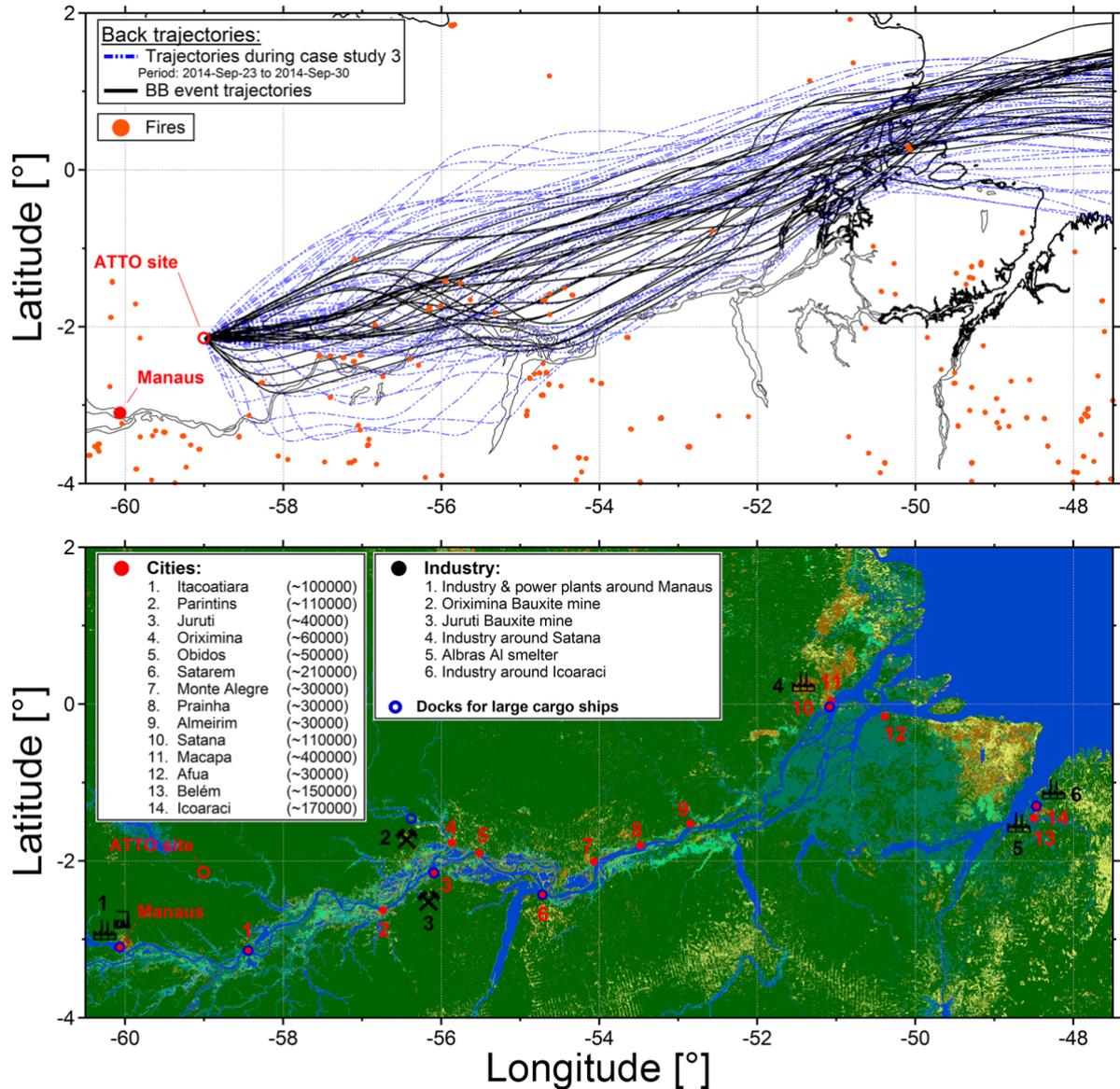


Figure S14. Composite maps illustrating conditions during *MPOL* case study (23 to 30 September 2014). Upper map combines backward trajectories ensemble for selected periods (HYSPLIT, NOAA-ARL, GDAS1, start height 200 m, trajectories started every 4 h) (Draxler and Hess, 1998) with satellite products for active fire counts (MODIS active fire detection extracted from MCD14ML distributed by NASA FIRMS) (Justice et al., 2011) during same period. Blue trajectories represent ensemble during entire *MPOL* period. Black trajectories highlight episodes with can be attributed to biomass burning plumes (24 Sep 2014 2200UT – 26 Sep 0800UT, 27 Sep 0000UT – 27 Sep 0600UT, 29 Sep 0000UT – 29 Sep 1000UT). Main trajectory track follows westerly direction and passes several (smaller fires) along Amazon River. Lower map shows GlobCover 2009 data (Arino et al., 2008) and larger cities (population given in parenthesis) as well as larger industrial infrastructure along the Amazon River as further potential pollution sources.

Table S1. Properties (position x_0 , integral number concentration N_{CN} , width σ) of Aitken and accumulation modes from double log-normal fits of the total particle size distribution for the PR_{BC} , PR_{CO} , and $PR_{BC \cup CO}$ cases as supplementary information to Table 2. The errors represent the uncertainty of the fit parameters. The error in $S_{cloud}(D_H, \kappa)$ is the experimentally derived error in S . For all double log-normal fits shown here, the fitted size range was limited to 50 - 350 nm since presence of particles in nucleation mode size range (<50 nm) interferes with fit of Aitken mode. All $N_{CN}(D)$ and $\kappa(S, D_a)$ size distributions specified here are plotted in Fig. S5. The following function has been used for the fitting of the experimentally obtained size distributions:

$$f(x) = y_0 + a_1 * e^{-\left(\frac{\ln\left(\frac{x}{x_1}\right)}{\sigma_1}\right)^2} + a_2 * e^{-\left(\frac{\ln\left(\frac{x}{x_2}\right)}{\sigma_2}\right)^2}$$

Based on this function and the data in this table, the characteristic size distributions in this work can be implemented into modelling studies.

Conditions	Mode or size range	N_{CN} [cm ⁻³]	κ	y_0 [cm ⁻³]	a_{mode} [cm ⁻³]	x_{mode} [nm]	σ_{mode}	R^2	D_H [nm]	$S_{cloud}(D_H, \kappa)$ [%]	
empirically pristine rain forest (PR) aerosol (PR_{BC})	Aitken	163 ± 5	0.13 ± 0.01	15 ± 2	416 ± 2	67 ± 1	0.47 ± 0.01	0.99	114 ± 2	0.23 ± 0.01	
	accumulation	83 ± 9	0.18 ± 0.01		195 ± 3	159 ± 1	0.44 ± 0.01				
empirically pristine rain forest (PR) aerosol (PR_{CO})	Aitken	158 ± 6	0.13 ± 0.02	22 ± 3	402 ± 3	67 ± 1	0.45 ± 0.01	0.99	109 ± 2	0.25 ± 0.01	
	accumulation	117 ± 1	0.19 ± 0.01		268 ± 4	156 ± 1	0.48 ± 0.01				
empirically pristine rain forest (PR) aerosol ($PR_{BC \cup CO}$)	Aitken	157 ± 5	0.13 ± 0.02	15 ± 3	417 ± 2	67 ± 1	0.45 ± 0.01	0.99	111 ± 2		
	accumulation	96 ± 4	0.18 ± 0.01		234 ± 3	157 ± 1	0.47 ± 0.01				
empirically pristine rain forest (PR) aerosol ($PR_{BC \cap CO}$ filter)	Aitken	162 ± 4	0.12 ± 0.01	15 ± 2	423 ± 2	69 ± 2	0.46 ± 0.01	0.99	103 ± 2	0.29 ± 0.01	
	accumulation	86 ± 8	0.18 ± 0.01		217 ± 3	157 ± 1	0.44 ± 0.01				
long-range transport (LRT) aerosol	Aitken	125 ± 12	0.18 ± 0.02	13 ± 2	250 ± 9	79 ± 3	0.60 ± 0.03	0.99	118 ± 2	0.15 ± 0.01	
	accumulation	313 ± 11	0.35 ± 0.04		770 ± 14	179 ± 2	0.52 ± 0.01				
biomass burning (BB) aerosol	Aitken	140 ± 29	0.14 ± 0.01	40 ± 23	688 ± 77	70 ± 1	0.20 ^b	0.99	--	--	
	accumulation	3439 ± 39	0.17 ± 0.02		7680 ± 49	167 ± 1	0.58 ± 0.01				
mixed pollution ($MPOL$) aerosol	$MPOL-LRT$	< 100 nm	0.14 ± 0.01	4 ± 7	1982 ± 10	135 ± 5	0.85 ± 0.01	0.99	--	--	
		> 100 nm									0.22 ± 0.03
	$MPOL-BB$	< 150 nm	2764 ± 84	0.10 ± 0.01	18 ± 21	4418 ± 30	113 ± 1	0.81 ± 0.01	0.99	--	--
		> 150 nm									

Table S2. Aerosol chemical composition from ACSM measurements at ATTO for *PR* filters as supplementary information to Table 3. ACSM data was available for time period from 01 Aug 2014 to 30 Sep 2016 and the averaged values are shown here as mean \pm std. The shown ACSM 3σ detection limits for 30 min averaging time were obtained from Ng et al. (2011). Note that M_{NH_4} ranged below the instruments detection limit in all *PR* cases, which makes the corresponding results unreliable. The predicted average aerosol hygroscopicity parameter, κ_p , was calculated according to the part 1 study (M. Pöhlker et al., 2016) and is shown as mean \pm se. The $\kappa(0.10 \%)$ results are shown with the experimentally derived error.

Time period	Mass concentrations M_{species} [$\mu\text{g m}^{-3}$] & (Mass fraction [%])						OA/SO_4^{2-}	κ_p	$\kappa(0.10 \%)$
	OA	NO_3^-	NH_4^+	SO_4^{2-}	Cl	BC_e			
empirically pristine rain forest (<i>PR</i>) aerosol (PR_{BC})	0.64 ± 0.01 (90)*	0.03 ± 0.03 (4)*	0.18 ± 0.01^a (--)*	0.04 ± 0.01 (6)*	$< (0.01 \pm 0.01)$ (--)*	$< (0.01 \pm 0.01)$ (--)*	~ 46	$0.16 \pm 0.01^*$	0.19 ± 0.05
empirically pristine rain forest (<i>PR</i>) aerosol (PR_{CO})	0.77 ± 0.01 (82)*	0.04 ± 0.01 (4)*	0.17 ± 0.01^a (--)*	0.09 ± 0.01 (9)*	0.01 ± 0.01 (1)*	0.03 ± 0.01 (3)*	~ 30	$0.19 \pm 0.01^*$	0.19 ± 0.04
empirically pristine rain forest (<i>PR</i>) aerosol ($PR_{\text{BC} \cup \text{CO}}$)	0.75 ± 0.01 (83)*	0.04 ± 0.01 (4)*	0.17 ± 0.01^a (--)*	0.08 ± 0.01 (9)*	0.01 ± 0.01 (1)*	0.02 ± 0.01 (2)*	~ 30	$0.18 \pm 0.03^*$	0.19 ± 0.04
Detection limits of ACSM & MAAP	0.15	0.01	0.28	0.02	0.01	0.01	--	--	--

^a Here, measured M_{NH_4} ranged below the ACSM detection limit and the shown values are questionable. Therefore, mass fractions and κ_p were calculated by omitting M_{NH_4} as outlined in Sect. 2.1. The corresponding results are marked by *.

Table S3. Error function (erf) fit parameters describing CCN efficiency spectra $N_{\text{CCN}}(S)/N_{\text{CN},10}$ vs. S as model input data representing PR filters as supplementary information to Table 4. A reference concentration of $N_{\text{CN},10}$ has been used in all cases. N_{CN} is shown as mean \pm se. The errors in S_{mode} and w_{mode} represent the uncertainty of the fit parameters (see Sect. 2.2). All CCN efficiency spectra specified here are plotted in Fig. S8.

Time period	$N_{\text{CN}} [\text{cm}^{-3}]$	erf fit	mode	a_{mode}	$S_{\text{mode}} [\%]$	w_{mode}	R^2
empirically pristine rain forest (PR) aerosol (PR_{BC})	255 ± 2	single	1	1	0.47 ± 0.01	1.69 ± 0.06	0.99
		double	1 2	1 0.19 ± 0.06	0.11 ± 0.01 0.65 ± 0.06	0.47 ± 0.26 1.18 ± 0.14	0.99
empirically pristine rain forest (PR) aerosol (PR_{CO})	292 ± 2	single	1	1	0.38 ± 0.01	1.79 ± 0.07	0.99
		double	1 2	1 0.28 ± 0.07	0.11 ± 0.01 0.62 ± 0.08	0.52 ± 0.22 1.17 ± 0.18	0.99
empirically pristine rain forest (PR) aerosol ($PR_{\text{BC} \cup \text{CO}}$)	276 ± 2	single	1	1	0.41 ± 0.01	1.71 ± 0.06	0.99
		double	1 2	1 0.22 ± 0.07	0.11 ± 0.01 0.60 ± 0.07	0.50 ± 0.27 1.21 ± 0.16	0.99

Table S4. Excess $N_{CCN}(S)$ to excess c_{CO} ratios, $\Delta N_{CCN}(S)/\Delta c_{CO}$, for the individual S levels during peak period of the biomass burning event in case study BB (17 - 23 August 2014, see Table 1). Ratios $\Delta N_{CCN}(S)/\Delta c_{CO}$ were obtained from bivariate regression fits (see Fig. S13b).

S	$\Delta N_{CCN}(S) / \Delta CO$	N	R^2
[%]	[$\text{cm}^{-3} \text{ppb}^{-1}$]	[cm^{-3}]	
0.11±0.01	6.7±0.5	-603 ±125	0.86
0.15±0.02	13.6±1.4	-1447 ±354	0.68
0.20±0.02	14.3±0.8	-1128 ±208	0.90
0.24±0.03	16.8±1.0	-1460 ±261	0.86
0.29±0.03	17.4±1.3	-1378 ±296	0.83
0.47±0.04	20.1±1.7	-1675 ±425	0.84
0.61±0.06	17.9±1.3	-1206 ±332	0.88
0.74±0.08	16.5±1.3	-933 ±329	0.88
0.92±0.11	18.1±1.4	-1265 ±355	0.85
1.10±0.08	17.5±1.3	-1096 ±328	0.87

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