Formation of hydroxyl radicals from photolysis of secondary organic aerosol material


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Abstract. This paper demonstrates that OH radicals are formed by photolysis of secondary organic aerosol (SOA) material formed by terpene ozonolysis. The SOA is collected on filters, dissolved in water containing a radical trap (benzoic acid), and then exposed to ultraviolet light in a photochemical reactor. The OH formation rates, which are similar for both α-pinene and limonene SOA, are measured from the formation rate of p-hydroxybenzoic acid as measured using offline HPLC analysis. To evaluate whether the OH is formed by photolysis of H₂O₂ or organic hydroperoxides (ROOH), the peroxide content of the SOA was measured using the horseradish peroxidase-dichlorofluorescein (HRP-DCF) assay, which was calibrated using H₂O₂. The OH formation rates from SOA are 5 times faster than from the photolysis of H₂O₂ solutions whose concentrations correspond to the peroxide content of the SOA solutions, assuming that the HRP-DCF signal arises from H₂O₂ alone. The higher rates of OH formation from SOA are likely due to ROOH photolysis, but we cannot rule out a contribution from secondary processes as well. This result is substantiated by photolysis experiments conducted with t-butyl hydroperoxide and cumene hydroperoxide which produce over 3 times more OH than photolysis of equivalent concentrations of H₂O₂. Relative to the peroxide level in the SOA and assuming that the peroxides drive most of the ultraviolet absorption, the quantum yield for OH generation from α-pinene SOA is 0.8 ± 0.4. This is the first demonstration of an efficient photolytic source of OH in SOA, one that may affect both cloud water and aerosol chemistry.

1 Introduction

Given the importance of secondary organic aerosol (SOA) to both climate and air quality, considerable attention has been given to studying the formation pathways and composition of SOA both in the lab and the field (Hallquist et al., 2009). However, less attention has been paid to its reactive properties and how the chemical nature of the particles may change as they transit through the atmosphere (Jimenez et al., 2009). If the particles become more hygroscopic, then they will be wet scavenged more easily. Similarly, their resultant health effects may transform if reactive functional groups are either produced or lost during processing.

One direction has been to address the multi-phase oxidation processes in which SOA participates, primarily via oxidation by gas-phase OH radicals (George and Abbatt, 2010). The conclusions from these studies are that heterogeneous exposure to OH leads to a more oxidized and hygroscopic aerosol, with small amounts of mass loss through fragmentation reactions occurring on a timescale of a few days of equivalent OH exposure in the atmosphere.

A second, less explored direction for SOA processing studies has been with respect to photochemical aging in the presence of ultraviolet light. In initial studies, many by Nizkorodov and co-workers, it has been shown that oxidative aging can also occur, leading to the formation of small molecules in the gas phase (e.g. HCHO, HCOOH, CO), rapid loss of condensed phase carbonyls, and some degree of mass loss (Walser et al., 2007; Mang et al., 2008; Bateman et al., 2011; Henry and Donahue, 2012; Epstein et al., 2014; Wong et al., 2014). The rates of this chemistry are dependent on environmental conditions, such as relative humidity (Wong et al., 2014). It is not clear from these experiments whether...
the aging occurs through only primary photochemical reactions or whether secondary processes also occur, initiated by radical production in a primary photochemical step. These photochemical aging experiments may be especially important for the change in optical properties of organic aerosol, especially for photobleaching of brown carbon species that absorb in the visible and near UV range (Sareen et al., 2013; Lee et al., 2014; Zhong and Jang, 2014; Zhao et al., 2015).

To further our understanding of this photochemical aging mechanism, it is necessary to evaluate the potential for oxidant production within tropospheric particles. Although it is known that gas-phase OH radicals collide with particles giving rise to oxidation, probably at the surface of the particle, the intrinsic sources of OH within a particle have not been well quantified. There are a number of routes that OH could form (Herrmann et al., 2010). The presence of soluble iron, for example, could lead to OH production through the Fenton reaction. Photochemical processes have been suggested as well, primarily through analogy to chemistry occurring in cloud water, such as the production of OH through the photolysis of efficient photochemical sources such as dissolved nitrate, nitrite, and hydrogen peroxide.

In this paper we measure the formation rate of OH radicals that occurs when SOA constituents are photolyzed, most likely through the photodissociation of organic hydroperoxide (ROOH). The presence of ROOH in SOA, especially ozonolysis SOA formed from terpenes, is likely (Ehn et al., 2014). They may form via Criegee biradicals, generated when ozone reacts with a carbon–carbon double bond, and also through H-abstractions that are part of chain oxidation mechanisms (Crounse et al., 2013; Ehn et al., 2014). Indeed, it is likely that the molecular structures of extremely low volatility organic compounds (ELVOCs) formed in the ozonolysis of α-pinene consist of numerous hydroperoxide functional groups mounted on the background of ring-opened functionalized α-pinene starting material. Secondary chemistry may also occur in aerosol subsequent to initial photolytic formation of radicals, producing additional photoactive ROOH species.

A number of past studies have quantitatively investigated the association of different peroxide species with SOA (Li et al., 2002; Docherty et al., 2005; Chen and Hopke, 2009a, b, 2010; Wang et al., 2011; Bateman et al., 2011; Mertes et al., 2012; Mutzel et al., 2013). In a study conducted by Li et al. (2002), the concentrations of H$_2$O$_2$ and ROOH formed during the ozonolysis of limonene were measured in both the gas and particle phases simultaneously. It was estimated that roughly 1 ppb of peroxide was generated from the reaction at limonene and ozone levels relevant to indoor conditions. A study by Docherty et al. (2015) measured the yields of total organic peroxides in SOA formed through the ozonolysis of α- and β-pinene, Δ-3-carene, and sabinene, reporting high yields of 47–85 % of the SOA mass. Chen and Hopke measured peroxides associated with particles formed using α-pinene, limonene, and linalool VOC precursors and measured their stability under different conditions (Chen and Hopke, 2009a, b, 2010); this was followed by the work of Wang et al. (2011) who also investigated peroxide formation and stability associated with SOA formed through the ozonation of α- and β-pinene and toluene precursors. Mertes et al. (2012) generated α-pinene SOA, measuring peroxide yields of 12–34 % of the SOA mass. A study by Bateman et al. (2011) measured the yields of peroxides from limonene SOA, reporting a value of 2 % in terms of the moles of SOA collected. The stability of the peroxides under photolytic conditions was also evaluated, finding that there was no significant change in peroxide levels following 14 h of photolysis, probably because of the formation of smaller peroxides during photo-processing.

The importance of a photochemical OH source in SOA is that the OH radical may lead to oxidation processes within aerosol particles and cloud water. In aerosol particles in particular, with small aqueous volumes, the partitioning properties of volatile species such as H$_2$O$_2$ indicate that little is expected to reside in the particles. Additionally, it is possible that other OH sources, such as nitrate ions, are not fully mixed with the SOA, especially when the particle has phase separated into inorganic-rich and organic-rich components. However, OH generated from SOA will be well mixed on a molecular scale with other SOA materials and so able to drive oxidative processes. A recent laboratory study has illustrated how OH generation within particles can lead to rapid oxidation of organic constituents (Daumit et al., 2014).

We describe experiments in which we photolyze SOA material generated in an environmental chamber after it is collected and then dissolved in water. We use an aqueous radical trap, benzoic acid (Klein et al., 1975; Anastasio and McGregor, 2001), to measure the OH production rate. To relate OH generation rates to the composition of the SOA, we apply a standard assay (horseradish peroxidase-dichlorofluorescein, HRP-DCF) to measure the peroxide content of the aerosol components (Keston and Brandt, 1965). Assuming that the peroxide signal arising from the HRP-DCF assay is due to the presence of H$_2$O$_2$, we compare the OH production rates from dissolved SOA to the rate from solutions with corresponding concentrations of H$_2$O$_2$. We find that the OH production rates from SOA are substantially higher than from the pure H$_2$O$_2$ solutions, implying that species other than hydrogen peroxide – likely organic hydroperoxide – are photolyzing into OH. To our knowledge, this is the first quantitative evaluation of the potential for SOA material to photolyze to form OH. As part of the study, we also perform detailed stability tests of the peroxides within SOA, as measured with the HRP-DCF assay, both in solution and on filter to better establish their environmental relevance.
2 Methods

2.1 SOA collection using an environmental chamber

An environmental chamber was used to collect SOA samples both to evaluate peroxide yields and for the photolysis experiments (see Fig. 1a). The chamber is a 1 m³ Teflon bag supported by a Teflon-coated frame. The bag is externally surrounded by metal panels, preventing exposure to outside light. The chamber was operated in continuous mode with all flows controlled by mass flow controllers. Ozone was generated by flowing purified air over a 185 nm mercury pen-ray lamp, with a flow rate of 6 standard L min⁻¹ (MFC 5, see Fig. 1a). A dilution flow of air of 7 standard L min⁻¹ (MFC 4) is mixed with the ozone flow prior to being introduced into the chamber. Limonene was introduced to the chamber using a 10 sccm flow of nitrogen through a headspace bubbl er chilled at 5 °C (MFC 1). α-Pinene was introduced to the chamber through a 12 sccm flow from a custom cylinder with a certified mixing ratio (320 ppm ± 20 % α-pinene in nitrogen, Air Liquide). The VOC flows were carried with a dilution flow of air of 500 sccm (MFC 2), meeting an additional dilution flow of 2.2 standard L min⁻¹ (MFC 3) before entering the chamber through a stainless steel port. The total flow rate through the chamber was held at 15 standard L min⁻¹ using a diaphragm pump (MFC 6). Aerosol samples were collected on supported PTFE filters (Zefluor, Pall Life Sciences, 47 mm diameter, 2.0 µM pore size) for 2 h, collecting an average of 185 and 900 µg of α-pinene and limonene SOA respectively.

2.2 SOA storage and extraction

Following collection, SOA samples were either extracted immediately or else stored for stability studies. To extract, filter samples were placed in a foil-covered Teflon bottle with 15 mL deionized water (18 mΩ, Millipore) and placed on a shake table for 15 min at 420 rpm. Following extraction, an aliquot of each sample was immediately prepared for analysis using the HRP-DCF assay or used for photochemistry experiments.

To evaluate in-solution stability, extracted solutions were covered with parafilm and foil and stored in the dark at room temperature. Those for on-filter stability testing were weighed and immediately placed in a sealed plastic filter holder (Analyslide petri dish, Pall Life Sciences). For the stability studies, samples were placed either in a dark cupboard at room temperature or in a sealed plastic bag and stored in a dark freezer at −20 °C. Frozen samples were brought to room temperature before being removed from the filter holder to prevent the condensation of contaminants onto the filter surface.

2.3 Horseradish peroxidase-dichlorofluorescein assay

A stock solution of 2′,7′-dichlorofluorescein diacetate (1 mM DCFHDA, C₂₄H₁₄C₂O₇, Sigma-Aldrich) was prepared in methanol and stored at −20 °C. The DCFHDA was converted to the hydrolyzed 2′,7′-dichlorofluorescein (DCFH) form by adding 1.0 mL DCFHDA to 4.0 mL sodium hydrox ide (0.01 M, NaOH) and allowing the hydrolysis to proceed for 30 min at room temperature. The fluorescing solution, referred to as HRP-DCF, was prepared by mixing 4.0 mL of DCFH with 7.0 mg of peroxidase from horseradish (HRP, Type I, Sigma-Aldrich), brought to a final volume of 100 mL with pH 7.2 phosphate buffer (7.35 mM KH₂PO₄, 17.6 mM Na₂HPO₄). The HRP-DCF solution was kept in amber bottles on ice until needed.

Hydrogen peroxide standards were freshly prepared for the calibration of each assay. A stock solution (1 mM H₂O₂,
made with H$_2$O$_2$ 30 % wt ACS reagent, Sigma Aldrich) was prepared in deionized water, used for standard dilutions ranging from 2 to 20 µM H$_2$O$_2$. Calibration standards were stored in amber bottles on ice until needed. The limit of detection of the HRP-DCF assay is 0.1 µM H$_2$O$_2$.

Calibration curves were also prepared using commercial sources of t-butyl hydroperoxide (Luperox TBH70X, 70 wt % in H$_2$O, Sigma Aldrich), cumene hydroperoxide (technical grade, 80 %, Sigma Aldrich), and di-t-butyl peroxide (Luperox DI, 98 %, Sigma Aldrich). The signal responses were respectively 97, 94, and 96 % lower than for hydrogen peroxide. Commercially available organic peroxides are those that are relatively stable. Thus, we conclude that the HRP-DCF response will be either due to dissolved H$_2$O$_2$ or to peroxides more unstable than those commercially available.

Assay samples were prepared by adding 250 µL of the sample to 2.25 mL HRP-DCF in Teflon-capped amber vials. Samples were briefly mixed before reacting at room temperature in the dark for 30 min. The reaction was quenched by putting vials on ice for 30 s before measuring the sample fluorescence. A spectrometer/fluorometer (SpectroVis Plus, Vernier) was operated in fluorescence mode with a 500 nm excitation light source, measuring sample spectra at 523.9 nm. Data were collected using Logger Pro software (version 3.8.2, Vernier).

2.4 Measurement of OH production rates

Photolysis of 50 mL volumes of SOA solutions was conducted in a photochemical reactor in which a 100 mL glass vessel was situated in the centre of an array of UV-B fluorescent lamps (see Fig. 1b). The spectrum of the lamps was measured with a spectral radiometer (StellaNet Inc.) and the magnitude of the flux (see Fig. 2) was calibrated by measuring the photoisomerization rate of 2-nitrobenzaldehyde (for details, see Zhao et al., 2015).

SOA solutions were prepared by mixing equal volumes of aqueous solutions of SOA and benzoic acid (BA), a known OH radical trap, resulting in final SOA and BA concentrations of 250 µM and 1.0 mM respectively where the SOA material is assumed to have an average effective molecular weight of 200 g mole$^{-1}$ (Bateman et al., 2011). The solutions were photolyzed for 1 h with 2.5 mL samples taken at 30 and 60 min for subsequent analysis. OH production rates are measured by the formation of para-hydroxybenzoic acid (PHBA), measured using HPLC, where all sample solutions were acidified immediately prior to injection with 2 drops of 1.0 mM H$_2$SO$_4$. A glass injection syringe was rinsed five times with deionized water and five times with sample solution. Sample solutions were analyzed using a 150 mm C-18 column with fixed wavelength detector at λ = 256 nm and detector range = 0.01. A 15 min, four-step gradient elution method was employed using mobile phase composed of acetonitrile and trifluoroacetic acid in deionized water.

PHBA forms in solution from OH reacting with benzoic acid with a yield of 20 % (Klein et al., 1975; Anastasio and McGregor, 2001). Based on experiments conducted with 0.25, 0.50, and 1.0 mM benzoic acid, it was determined that 1.0 mM solutions trap all the OH generated. In particular, more OH was trapped in SOA photolysis experiments conducted with 0.50 mM BA solutions than with 0.25 mM solutions, but the amount trapped in 0.50 and 1.0 mM runs was the same within experimental error. Overall, the detection limit for OH production rate within the SOA solutions was on the order of 1 × 10$^{-10}$ M s$^{-1}$, as determined largely by background amounts of PHBA present in the solutions without illumination.

Photolysis experiments were also performed with 12.5 µM solutions of commercial organic hydroperoxides, namely t-butyl hydroperoxide and cumene hydroperoxide, and hydrogen peroxide.

3 Results and discussion

3.1 SOA peroxide yields and thermal stability

A series of experiments was conducted with the HRP-DCF assay both to determine how best to handle SOA samples to preserve their peroxide content and also to gain some information on the nature of the peroxides present in the SOA, especially when dissolved in water. In particular, HRP-DCF peroxide yields (see Table 1) from fresh SOA samples were measured immediately following collection and the 15 min aqueous extraction period. The yields have been reported in three ways: (1) mole percentage, moles of peroxides/moles of SOA × 100 %, where the molecular weight of SOA is assumed to be 200 g mole$^{-1}$ (Bateman et al., 2011); (2) mass percentage, mass of peroxides/mass of SOA collected × 100 %, where the molecular weight of peroxides is assumed to be 34 g mole$^{-1}$; and (3) normalized yield, moles of peroxides/mass of SOA collected, where the peroxides are assumed to be H$_2$O$_2$.
Peroxide yields were also measured using SOA generated in a flow tube (see Supplement, Fig. S1) with mass loadings at least 10 times higher than those generated in the chamber. These results demonstrate that mass loading does not have a large effect on peroxide yields in SOA, at least over the range of conditions explored. For example, the fresh yields from chamber SOA were 3.1 ± 0.5 % mole for α-pinene and 5.1 ± 0.1 % mole for limonene, very similar to those from SOA collected using the flow tube: 4.2 ± 0.6 % mole and 3.8 ± 0.7 % mole respectively. Uncertainties quoted reflect variability from between three and seven replicates, whereas we estimate absolute uncertainties are on the order of ±20 %. The HRP-DCF peroxide yields from α-pinene and limonene SOA only comprise a few percent of the total moles of SOA material. These values are comparable to the results reported by Chen and Hopke, who also used the HRP-DCF assay to measure peroxides from SOA formed through α-pinene and limonene ozonolysis (Chen and Hopke, 2009a, 2010). The studies by both Docherty et al. (2005) and Mertes et al. (2012) used the iodide technique to quantify their peroxide species. This technique is sensitive to organic peroxides (i.e. ROOH and ROOR) and therefore the yields reported by these studies are considered to be the total peroxide content of the SOA. The study by Bateman et al. (2011) also used the iodide method to quantify peroxides, using H₂O₂ to calibrate the fluorescence response. The other studies using the iodide method calibrated the assay using benzoyl peroxide, an organic peroxide that may have a different sensitivity to the assay than hydrogen peroxide. Finally, the study by Wang et al. (2011) reported yields that were somewhat lower than the α-pinene yields in this study. This study also used HRP as their catalyst; however, they used an HPLC technique to speculate the peroxides, and para-hydroxyphenylacetic acid was used in the place of dichlorofluorescein.

As noted in the Methods section, the sensitivity of the HRP-DCF assay heavily favours H₂O₂ although a weak sensitivity to commercial organic hydroperoxides exists. A control experiment was conducted in which a very high gas-phase concentration of H₂O₂ (created by bubbling 250 sccm N₂ through 30 % H₂O₂ solution) was passed through a Teflon filter for 1 h. Upon extraction in water and HRP-DCF analysis, it was found that no measurable amount of H₂O₂ had adhered to the filter. This indicates that the peroxides measured in this work were not in the form of H₂O₂ when present in the aerosol particles, as expected from the high volatility of H₂O₂. Rather, the signal is most likely due to ROOH in the particle. These molecules may have a direct response in the HRP-DCF assay, higher than the stable commercial organic hydroperoxides. Also, it is possible that some of the signal arises from H₂O₂ when SOA material is put into solution. For example, a general class of compounds, the α-hydroxyhydroperoxides, exist in equilibrium in water with hydrogen peroxide and organics containing carbonyl functional groups, especially aldehydes (Zhao et al., 2013). These species can be formed through the hydrolysis reaction of Criegee intermediates. Similar conclusions have been made previously (Wang et al., 2011).

Support for the formation of some H₂O₂ in solution comes from a set of stability experiments conducted both in solution and on the filter, using SOA generated in the flow tube. As shown in the Supplement (see Fig. S2), the stability in aqueous solutions of α-pinene and limonene SOA had very similar trends following extraction. Within the first 48 h after extraction, the amount of peroxide in solution increased by approximately 50 % of its original yield. After 48 h, the peroxide slowly decayed. A control experiment was performed in which the Teflon filter was removed from the aqueous solution following the initial 15 min extraction period. In this sample, the same increase in peroxide yields was observed within the first 48 h following extraction. If the peroxides in the particle begin as organic hydroperoxides, the increase in yield may be due to a gradual decomposition reaction to form H₂O₂ within the first 48 h. Although the total peroxide content in the extract may not be changing, a larger fraction would be present in the H₂O₂ form. Since the HRP-DCF assay is very sensitive to H₂O₂, this could account for the in-

<table>
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<tr>
<th>α-Pinene</th>
<th>Detection method</th>
<th>Yield (% mole)</th>
<th>Yield (% mass)</th>
<th>Normalized (mole μg⁻¹)</th>
</tr>
</thead>
<tbody>
<tr>
<td>This study – chamber</td>
<td>HRP-DCF</td>
<td>3.1 ± 0.5</td>
<td>0.5 ± 0.1</td>
<td>(1.6 ± 0.3) × 10⁻¹⁰</td>
</tr>
<tr>
<td>Docherty et al. (2005)</td>
<td>Iodide</td>
<td>–</td>
<td>47 ± 12</td>
<td>–</td>
</tr>
<tr>
<td>Chen and Hopke (2009a)</td>
<td>HRP-DCF</td>
<td>–</td>
<td>–</td>
<td>(1.8 ± 0.8) × 10⁻¹⁰</td>
</tr>
<tr>
<td>Wang et al. (2011)</td>
<td>HRP-PHOPAA*</td>
<td>–</td>
<td>–</td>
<td>(2.7 ± 1.1) × 10⁻¹¹</td>
</tr>
<tr>
<td>Mertes et al. (2012)</td>
<td>Iodide</td>
<td>–</td>
<td>34 ± 4</td>
<td>–</td>
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</table>

Limonene

<table>
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<tr>
<th>Detection method</th>
<th>Yield (% mole)</th>
<th>Yield (% mass)</th>
<th>Normalized (mole μg⁻¹)</th>
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<td>This study – chamber</td>
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<td>5.1 ± 0.1</td>
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<td>–</td>
<td>–</td>
</tr>
<tr>
<td>Bateman et al. (2011)</td>
<td>Iodide</td>
<td>2</td>
<td>–</td>
</tr>
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*PHOPAA is para-hydroxyphenylacetic acid
The rates of PHBA formation from the 250 µM 3.2 OH generation rates from photolysis (Wang et al., 2011). This is in accord with previous work (Chen and Hopke, 2010; Zhao et al., 2013). We note in general that methods for the accurate speciation and quantification of complex ROOH species are poorly developed, and limit some of the conclusions that can be made in this work and in the field in general.

The peroxides present in the particles have potential to influence photochemical pathways, as will be discussed below, but also to affect human health. SOA can be either inhaled directly or can deposit on indoor surfaces where it can become part of the semi-volatile organic matter that is known to be ubiquitously present (Liu et al., 2003). For the airborne particles, peroxides associated with the SOA, as one component of the general class of molecules referred to as reactive.

4 Conclusions and environmental implications

4.1 SOA peroxides

Although total peroxide yields have been reported previously there is considerable uncertainty in their interpretation as described above, dependent in part on the methods of analysis (see Table 1). The assays that have been used are non-specific, with the iodide approach sensitive to both ROOH and ROOR species and the HRP-DCF method sensitive to H₂O₂ and probably some ROOH species. As described above, the peroxide stability studies indicate that the HRP-DCF peroxide signal is stable when filters are stored frozen, likely because there is no decomposition and subsequent loss of volatile species that are HRP-DCF active. Also, the solution studies show that a more active species, likely H₂O₂, is slowly formed in solution within the first 24 h perhaps from the decomposition of soluble species such as α-hydroxyhydroperoxides. These molecules are known to decompose to form hydrogen peroxide and organics with carbonyl functional groups (Wang et al., 2011; Zhao et al., 2013).
oxygen species (ROS), have the potential to cause oxidative stress upon inhalation. Based on our work, it is expected that the peroxides exist as organic hydroperoxides in the particles – and not as H₂O₂ – thus allowing them to be carried deep into the respiratory system. Being volatile and soluble, it is unlikely that gas-phase H₂O₂ can pass long distances into the respiratory pathways. Once the particles deposit and dissolve in the lung fluid, the organic hydroperoxides may decompose to form H₂O₂, a potent ROS constituent.

4.2 Photochemical generation of OH

The primary result from this work is that photolysis of SOA material generates OH in solution, likely from ROOH species. In particular, the quantum yield for gas-phase t-butyl hydroperoxide is unity (Baasandorj et al., 2010). Also, the effective quantum yield for OH formation from H₂O₂ in solution is also close to unity (Herrmann et al., 2010; Goldstein et al., 2007).

For this work, to calculate the overall quantum yield of OH from α-pinene SOA peroxides we first calculate the first-order rate constant (k_OH) describing OH formation:

\[
[\text{OH}]_t = [\text{Peroxide}]_0 (1 - \exp(-k_{\text{OH}}t)),
\]

where [OH]₀ refers to the time-dependent concentration of OH and [Peroxide]₀ represents the maximum concentration of OH that will form from peroxide photolysis at infinite time, assuming that one OH radical can form from one peroxide molecule. To make this calculation, we convert the PHBA formation rates in Fig. 3 (i.e. the slopes of the lines of best fit) into OH formation rates assuming a yield of 20% of PHBA from reaction of OH with benzoic acid (Klein et al., 1975; Anastasio and McGregor, 2001). The results are given in Table 2. Using Eq. (1), the rate constant for OH production from α-pinene SOA photolysis is thus calculated to be (1.4 ± 0.4) x 10⁻⁴ s⁻¹, where the uncertainty arises from estimated absolute uncertainties in the measured OH production rate and the peroxide concentration of the α-pinene SOA solution.

We then express this photolysis rate constant, k_OH, in terms of the wavelength (λ)-dependent light intensity in the reactor (I(λ)), the SOA absorption cross section (σ(λ)), and the quantum yield (φ(λ)):

\[
k_{\text{OH}} = \int I(\lambda) \sigma(\lambda) \phi(\lambda) d(\lambda),
\]

where we have previously published the molecular absorption cross section for α-pinene SOA generated in the same environmental chamber (Wong et al., 2014). Note that by doing so we are referencing the quantum yield to the total concentration of peroxides in solution (i.e. assuming that the OH is arising from the peroxides) and we are assuming that the absorption cross section measured for SOA is determined by these peroxides in the pertinent wavelength range. While the absorption in the long-wavelength tail is likely due to other species, peroxides are likely to contribute to some degree to the short wavelength absorption. We do not know the degree to which other functional groups, such as carboxyls, also contribute.

By integrating this expression and equating it to the experimental k_OH value (i.e. (1.4 ± 0.4) x 10⁻⁴ s⁻¹), we calculate the peroxide quantum yield for aqueous α-pinene SOA to be 0.8 ± 0.4. This value represents an average effective quantum yield, assumed to be wavelength independent, for all routes to OH formation in solution. Its large value close to unity provides support that SOA peroxides are indeed the source of the OH measured. Using the same approach and literature values for the absorption cross section (Finlayson-Pitts and Pitts, 2000), we calculate the quantum yield for OH production from H₂O₂ in our experiments to be 1.1 ± 0.4, in agreement within experimental uncertainties with values in the literature (Goldstein et al., 2007; Herrmann et al., 2010).

We cannot determine from these data alone whether the OH is formed only in a primary step, as from the photolysis of a ROOH functional group that is part of the original SOA material, or whether secondary chemistry also contributes. For example, subsequent chemistry might involve the following:

ROOH + hν \rightarrow RO + OH \quad \text{(R1)}
RO \rightarrow R’ + R’HCO \quad \text{(R2)}
R’ + O₂ \rightarrow R’O₂ \quad \text{(R3)}
R’O₂ + HO₂ \rightarrow R’OOH + O₂ \quad \text{(R4)}
R’OOH + hν \rightarrow R’O + OH. \quad \text{(R5)}

The HO₂ that is needed in Reaction (R4) can be formed by photolysis of aldehydes that are part of the SOA material or from those generated in Reaction (R2),

R’HCO + hν \rightarrow R'' + HCO \quad \text{(R6)}
HCO + O₂ \rightarrow HO₂ + CO, \quad \text{(R7)}

and potentially from the oxidation of benzoic acid. The HO₂ can then go on to form H₂O₂ (Hullar and Anastasio, 2011),

HO₂ + HO₂ \rightarrow H₂O₂ + O₂, \quad \text{(R8)}
which may generate OH upon photolysis.

An indication that secondary chemistry may be occurring is that we would have expected the primary OH production rate to start to level off at 1 h reaction time due to the consumption of hydroperoxide species for the measured initial rate constant of $1.4 \times 10^{-4} \text{s}^{-1}$, if primary photolysis is the only production route of OH. The fact that the production rate remains linear with time may indicate a secondary source of OH. Catalytic formation of ELVOC-like compounds during the OH oxidation may represent a sustained source of hydroperoxides in addition to the mechanisms outlined above.

There are different approaches to assess the potential impact that SOA materials may have on condensed phase OH production rates in the atmosphere. To start, it is important to distinguish between cloud water and aerosol environments. For the very high liquid water volumes present in cloud water, highly soluble species that are also OH sources are present in reasonably high amounts. For example, H$_2$O$_2$ is present through Henry’s Law solubility at the 10 to 100 µM level for ambient gas-phase mixing ratios of 0.1 to 1 ppbv (Sakugawa et al., 1990). The corresponding concentration of SOA materials is not straightforward to estimate. Cloud water dissolved organic carbon (DOC) amounts range from a few mg CL$^{-1}$ up to many tens of mg CL$^{-1}$ in polluted environments (Herckes et al., 2013). Taking 10 mg CL$^{-1}$ as a representative value, an OM/OC ratio of 2, and assuming a molecular weight of 100 g mole$^{-1}$ for dissolved species (i.e. a combination of low molecular soluble species and higher molecular weight humic-acid-like molecules), this corresponds to a concentration of DOC species on the order of 200 µM. Only 10 to 50 % of DOC species has been identified at the molecular level as small soluble molecules (Herckes et al., 2013), making it possible that dissolved organics from SOA and other sources constitute the remaining fraction, i.e. roughly on the order of 100 µM. Thus, the SOA concentrations in cloud water may be comparable to those of H$_2$O$_2$ or even higher. In terms of OH-generating efficiency, Table 2 illustrates that SOA, on a per molar basis, is about 5 times less efficient than H$_2$O$_2$. However, the lower efficiency of OH production from SOA as compared to H$_2$O$_2$ may be offset by higher dissolved concentrations, making the SOA an important OH source in cloud water. Confirmation of this hypothesis requires better speciation of organic materials in cloud water that arise from SOA.

Following this line of thought forward to aerosol particles, the concentrations of dissolved species such as H$_2$O$_2$ will not be significantly higher than in cloud water. However, the concentrations of SOA materials will be very much larger, approaching molar values, which will make the SOA much more important as a photolytic OH source.

From another perspective, we can compare our OH photolytic formation rates to the formation rates of OH in cloud and aerosol that have been calculated based on known concentrations of constituents and known photochemistry and taking into account mass transfer from the gas phase (Arakaki et al., 2013). Formation rates vary widely from values of $10^{-10} \text{M s}^{-1}$ in rainwater to $10^{-9} \text{M s}^{-1}$ in cloud and fog to $10^{-7}$ to $10^{-6} \text{M s}^{-1}$ in marine aerosol particles. By comparison, we observed OH production rates of $10^{-9} \text{M s}^{-1}$ for our 250 µM SOA solutions. Given that 250 µM is not an inappropriate DOC concentration for cloud water conditions, especially if highly polluted (Herckes et al., 2013), this shows that the source has the potential to compete with more conventional OH sources, such as H$_2$O$_2$. The formation rate would be very much higher under aerosol water conditions given the high concentrations of SOA in such particles. We note that mass transfer of OH from the gas phase is also important for the overall OH flux and can compete with these condensed-phase formation pathways.

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